Analytical Solutions for Spontaneous Imbibition: Fractional-Flow Theory and Experimental Analysis

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Summary
We present analytical solutions for capillary-controlled displacement in one dimension by use of fractional-flow theory. We show how to construct solutions with a spreadsheet that can be used for the analysis of experiments as well as matrix-block-scale recovery in field settings. The solutions can be understood as the capillary analog to the classical Buckley-Leverett solution (Buckley and Leverett 1942) for viscous-dominated flow, and are valid for cocurrent and countercurrent spontaneous imbibition (SI), as well as for arbitrary capillary pressure and relative permeability curves. They can be used to study the influence of wettability, predicting saturation profiles and production rates characteristic for water-wet and mixed-wet conditions. We compare our results with in-situ measurements of saturation profiles for SI in a water-wet medium. We show that the characteristic shape of the saturation profile is consistent with the expected form of the relative permeabilities. We discuss how measurements of imbibition profiles, in combination with other measurements, could be used to determine relative permeability and capillary pressure.

Introduction
SI has received considerable attention in the literature because it is one of the key production mechanisms in fractured reservoirs that host the majority of the world’s remaining conventional oil. SI can occur in countercurrent and cocurrent mode. In countercurrent imbibition, water and oil both flow through the inlet of a core or matrix block but in opposite directions, whereas in cocurrent imbibition, oil and water flow in the same direction and the oil leaves through the other end. Both cocurrent and countercurrent SI are important in fractured reservoirs. If the matrix blocks are not fully surrounded by water, countercurrent imbibition prevails. Otherwise, cocurrent imbibition is dominant (Bourbiaux and Kalaydijian 1990; Pooladi-Darvish and Firoozabadi 2000). The difference between these mechanisms is crucial because experiments and numerical studies alike show that recovery rates and displacement efficiency are significantly higher in the case of cocurrent imbibition (Bourbiaux and Kalaydijian 1990; Pooladi-Darvish and Firoozabadi 2000; Chen et al. 2003). Therefore, any complete analytical description of SI must not only be capable of accounting for all relevant petrophysical properties, such as wetting and fluid properties, but must also be able to distinguish between cocurrent and countercurrent SI.

Although an analytical solution for purely advective flow has been known for decades—the classical Buckley-Leverett solution (Buckley and Leverett 1942)—its analytical counterpart for capillary-driven flow (SI) is less well-established in the literature (Schmid et al. 2011; Schmid and Geiger 2012, 2013). Given the outstanding practical importance of SI and the usefulness of analytical solutions, many authors have tried to obtain solutions for this situation. However, because the underlying strong nonlinearities pose significant mathematical difficulties, most analytical solutions made additional assumptions, such as specific relative permeability curves (Vortos and Fokas 1983), a linear capillary pressure-saturation relationship (Cil and Reis 1996), or steady-state flow (Kashchiev and Firoozabadi 2002). Barenblatt et al. (1990) provides a review of the traditional mathematical treatment of SI.

McWhorter and Sunada (1990) reported analytical solutions for cocurrent and countercurrent imbibition where no such additional assumptions were made. However, their mathematical formalism required that the boundary conditions for the water inflow be stated in a specific way, which led all subsequent authors—including McWhorter and Sunada (1992)—to believe that the solution obtained was yet another specific case, and the search for a general solution continued (Kashchiev and Firoozabadi 2002). Schmid et al. (2011) recently proved that the situation described by McWhorter and Sunada (1990) is a general, closed-form solution for SI, and showed that formally the solution might be understood as the Buckley-Leverett analog of SI. Later Schmid and Geiger (2012) showed that on the basis of this closed-form solution, the general nondimensional time or “scaling group” for SI can be obtained. Contrary to all other scaling equations, their scaling group does not require fitting parameters, and independent of rock and fluid type reliably scales water-wet (Schmid and Geiger 2012) and mixed-wet (Schmid and Geiger 2013) experiments. The closed-form solutions for SI, however, can look mathematically challenging because they require the calculation of two implicit integrals. This mathematical difficulty can easily obscure the simple physical meaning of the solution and make its presentation appear much more difficult than the Buckley-Leverett solution (Mason and Morrow 2013). The aim of this paper is therefore to give a novel introduction to the analytical solution for SI by use of physical reasoning, simplifying the mathematical formalism presented previously (McWhorter and Sunada 1990; Schmid et al. 2011) in a way that makes it accessible to anyone with some understanding of the Buckley-Leverett method and fractional-flow theory. We also provide a novel but simple way of calculating the solution in a spreadsheet. We then compare with measured saturation profiles for cocurrent SI and show how the analytical solutions can be used to constrain relative permeability and capillary pressure, as a complement to other types of measurement.

The paper is structured as follows. We start with introducing the equations for linear, incompressible two-phase flow. We show how, on physical grounds, an analytical solution for cocurrent and countercurrent SI can be derived in a way similar to the Buckley-Leverett theory, and how it can easily be computed with a simple spreadsheet. In the example section, we illustrate how to obtain the solution for water-wet and mixed-wet rocks. The examples also demonstrate how the analytical solutions predict the difference between cocurrent and countercurrent SI, and the influence of wetting behavior on displacement efficiency and recovery rates. In the application section, we compare our solutions with measured saturation profiles and use the theory to constrain the relative permeabilities and capillary pressures. The use of SI solutions to match to saturation measurements during SI, combined with other measurements to determine relative permeability and capillary pressure, is discussed.

The spreadsheets can be downloaded from https://www.imperial.ac.uk/engineering/departments/earth-science/research/research-groups/perm/research/pore-scale-modelling/software/ and http://carbonates.hw.ac.uk/software.html.
Two-Phase-Flow Equations

We consider 1D flow of two immiscible, incompressible phases through a homogeneous porous medium. Conservation of mass (strictly, volume) can be written as

\[ \phi \frac{\partial S_w}{\partial t} + \frac{\partial q_w}{\partial x} = 0, \]  

where \( \phi \) is the porosity; \( S_w \) and \( S_o \) are the water and oil saturation, respectively; and \( q_w \) is the water velocity (Dake 1983; Helming 1997):

\[ q_w = \frac{\lambda_w}{\mu_w} + \frac{\lambda_w \phi}{\mu_o} \frac{\partial P}{\partial x} + \frac{K}{\mu_o} \phi (P_o - \phi P_w), \]  

where \( \lambda_w \), \( \lambda_o \), \( w \) is the mobility of phase \( a \), \( w \); \( \lambda_w = \lambda_o + \lambda_w \); \( \lambda_a \) is the total mobility; \( \phi = q_o + q_w \) is the total velocity of water and oil; \( K \) is the absolute permeability; \( P \) is the capillary pressure defined as the difference between the oil and water pressure, or \( P_o - P_w \); \( \mu_a \) is the fluid viscosity; and \( g \) is the gravitational acceleration in the flow direction. The equation for \( q_w \) shows that water velocity is a result of the total velocity, a gradient in capillary pressure and gravity, as represented by the first, second, and third terms of the right-hand side of Eq. 2, respectively.

Analytical Solutions for Capillary-Dominated Cases

If we ignore capillary forces (the second term in Eq. 2 is set to zero), the Buckley-Leverett solution can be constructed. We will not repeat the analysis here because it is standard in the literature (Buckley and Leverett 1942; Dake 1983). The key point for our analysis is that, for a smooth variation in saturation (that is, ignoring the shock), we can write

\[ x(S_w, t) = \frac{q_o}{F_w}(S_w(t)), \]  

for a fixed injection rate \( q_o \). The total flow rate \( q_o \) is a free parameter; it is imposed as a boundary condition and its value scales the speed with which the water moves. If we ignore gravitational forces, the fractional flow of the water phase is defined as \( f_w = \lambda_w / \lambda_a \).

The Buckley-Leverett Analog for Capillary-Dominated Flow.

We now consider solutions to Eq. 1, including capillary pressure but ignoring gravitational forces (the third term in Eq. 2). In both cases, there is no forced injection of water (the wetting phase), but water spontaneously enters the porous medium under the action of capillary forces only. We also neglect nonequilibrium effects, which have been observed in the literature (Barenblatt et al. 1990; Le Guen and Kovscek 2006) because the model encapsulated in Eqs. 1 and 2 allows us to match our experimental data, as will be shown later.

The case we consider first is cocurrent flow, where the immiscible phase (water) enters the inlet and the displaced phase (oil) escapes from the opposite end. Oil and water flow in the same direction. In this case, there is a finite total velocity. The water velocity, from Eq. 2, is

\[ q_w = f_w(S_o)q_o + \frac{\lambda_w \phi}{\mu_o} \frac{\partial P}{\partial x} + \frac{K}{\mu_o} \phi (P_o - \phi P_w), \]  

For the countercurrent case, water invades a closed system, and therefore oil leaves in the opposite direction to the water. In this case \( q_o = -q_w \), which yields \( q_o = 0 \) and leads to

\[ q_o = -K \frac{\lambda_w \phi}{\mu_o} \frac{\partial P}{\partial x} + \frac{K}{\mu_o} \phi (P_o - \phi P_w). \]  

Notice the extra term for \( q_o \) in the cocurrent case. This implies that even if all petrophysical properties are the same, \( q_w \) in the cocurrent and countercurrent cases will be different. Because \( f_w(S_o)q_o \) is the injection rate, \( q_w \) will be larger than the inflow for the cocurrent case. This is indeed observed in experiments and numerical tests (Bourbiaux and Kalaydjian 1990; Pooiadi-Darvish and Firouozabadi 2000; Chen et al. 2003) that show that production rates from cocurrent imbibition are always larger than those from countercurrent SI on the same rock. \( f_w(S_o)q_o \) is known as the “viscous-coupling” term. Obviously, if \( f_w \) is small compared with the capillary term over the saturation range that is relevant for SI, then the viscous-coupling term will be small, and the difference between cocurrent and countercurrent SI becomes negligible. This is the case for mixed-wet systems or where the displaced phase has a low viscosity, such as air, which we will discuss further.

Let us return to finding an analytical solution for SI. From Eqs. 1 and 2, the conservation equation for \( S_w \) can be reduced to a nonlinear diffusion equation:

\[ \phi \frac{\partial S_w}{\partial t} = \frac{\partial}{\partial x} \left[ D(S_w) \frac{\partial S_w}{\partial x} \right], \]  

where \( D(S_w) \) is the nonlinear capillary-diffusion coefficient, with units of length squared over time (Barenblatt et al. 1990):

\[ D(S_w) = -K \frac{\lambda_w \phi}{\mu_o} \frac{\partial P}{\partial x}. \]  

Note that \( D(S_w) \) is zero if \( S_w = S_{wir} \). The implication of this, as we show later, is that the saturation gradient will diverge at the leading edge of the imbibition front, where \( S_w = S_{wir} \). We will find an analytical solution for Eq. 6 together with the boundary and initial conditions:

\[ S_w(x, t) = S_{wirs}, \]  

The injection condition is the maximum water saturation \( S_{wirs} \) possible under SI, and is the value where the capillary pressure is zero, or \( P_c(S_{wirs}) = 0 \). For weakly water-wet and mixed-wet media, \( S_{wirs} \) may be less than the maximum water saturation achievable for forced displacement, where the capillary pressure may be negative. The initial condition is a uniform water saturation, which we assume is the residual or connate value; that is, the water has zero relative permeability. The approach we follow is derived from the theory of nonlinear dispersive partial-differential equations (Tao 2006). The formal derivation has been developed previously (McWhorter and Sunada 1990; Schmid et al. 2011).

We motivate the solution by considering measurements of saturation profiles for SI in Kettoll limestone shown in Fig. 1. Similar profiles have been measured by Le Guen and Kovscek (2006) and Zhou et al. (2002). Fig. 2 shows the average saturation as a function of distance from the inlet divided by the square root of the time since the start of the displacement for the experiment shown in Fig. 1 and for an experiment in diatomite (Zhou et al. 2002). Fig. 3 shows the capillary pressure and relative permeabilities used in the analytical solution, which we discuss later. To a good approximation, the measurements fall onto one curve for different times. Whenever two different times and locations satisfy \( x_1 / t_1 = x_2 / t_2 \), then also the saturation values at \( x_1 \) and \( x_2 \) are the same at the respective times \( t_1 \) and \( t_2 \). Although the distance traveled by any saturation level is proportional to \( \sqrt{t} \), different saturation levels are transported with different overall velocity, and not simultaneously. Therefore, the distance traveled must also be a function of saturation. This is similar to the Buckley-Leverett theory, but the key difference is that the diffusive nature of Eq. 6 gives rise to a \( 1/\sqrt{t} \) dependence rather than a linear dependence in time. Why do we see this diffusive behavior? If we have a displacement controlled entirely by capillary forces, then the driving force is simply the capillary pressure, which remains fixed for a given saturation. However, from Darcy’s law, the flow rate is proportional to the pressure gradient. Hence, if the saturation has moved a distance \( x \), then the pressure gradient is proportional to \( 1/x \). The flow rate gives the speed of movement \( dx/dt \) and hence a simple integration gives \( x^2 \approx t \) or \( x \approx \sqrt{t} \).

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By use of this physical observation, we introduce the similarity variable \( \omega = x \sqrt{t} \) and attempt to find solutions for \( S_w \) that are a function of \( \omega \) only (Barenblatt et al. 1990):

\[
\omega = \frac{x}{\sqrt{t}}, \quad S_w = S_w(\omega).
\]

Because the distance traveled by a given saturation value is proportional to \( \sqrt{t} \), the water volume imbibed must be proportional to \( \sqrt{t} \). Therefore, \( q_w(x = 0, t) \) scales as \( 1/\sqrt{t} \). This means that we can find a proportionality constant \( C \) in m/\( \sqrt{s} \) units (\( C^2 \) has the units of a diffusion coefficient, as does \( D \) in Eq. 7), for which we can write

\[
q_w(x = 0, t) = \frac{C}{\sqrt{t}},
\]

\[
Q_w(t) = \int_{t=0}^{t} q_w(x = 0, t) \, dt = 2C \sqrt{t}.
\]

The magnitude of \( C \) determines how fast the water spontaneously imbibes, and therefore depends on the properties of the rock and fluids, such as the wetting properties, fluid viscosities, and relative and absolute permeabilities. Also, it is clear that because \( q_w \) is different for cocurrent and countercurrent flow, \( C \) must be different for these two cases as well. Once we know how to calculate \( C \), we have a simple way of quantifying how rock and fluid properties influence a rock’s ability to imbibe in cocurrent and countercurrent mode (this is discussed later with our examples).

The key new step to finding a general solution is to write the solution in terms of the derivative of a fractional-flow function \( F \) (Schmid et al. 2011):

\[
\omega(S_w) = \frac{C}{\bar{q}} F(S_w). \tag{11}
\]

Eq. 11 is formally identical to the Buckley-Leverett solution, which is Eq. 3. We therefore consider \( F \) to be the fractional-flow function for the case of capillary-driven flow, and we hypothesize that \( F \) is given by the ratio of the water flow as a function of saturation to its maximum value:

\[
F = \frac{q_w}{q_w(S_{w,max})}. \tag{12}
\]

\[
\begin{align*}
\text{Fig. 1—Measured SI profiles for a cocurrent experiment. The saturation front moves into the core with a time dependence of } \sqrt{t}. \\
\text{This observation can be used for constructing an analytical solution. In the experiment, water imbibes from the base of a Ketton-} \\
limestone-core sample, whereas air (the displaced phase in this case, as opposed to oil) escapes from the top. The saturation profiles are measured by use of X-ray computed tomography scanning.}
\end{align*}
\]

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\[
\begin{align*}
\text{Fig. 2—The saturation profiles from Fig. 1, showing the average saturation along the flow direction for Ketton limestone in (a) } \\
\text{cocurrent flow and (b) for diatomite for countercurrent flow obtained from Zhou et al. (2002). The profiles are plotted as functions } \\
\text{of } x \sqrt{t}, \text{where } x \text{ is the distance moved in the flow direction and } t \text{ is the time since the beginning of the experiment. To a reasonable approximation, profiles measured at different times fall onto one universal curve. This indicates that the solution for capillary-} \\
\text{dominated displacement in one dimension can be written as a function of } x \sqrt{t} \text{ only. We use this insight to construct analytical solutions. Solutions that match the experiments are shown by the solid black lines. The multiphase-flow properties used for this match are shown in Table 1, while the relative permeabilities and capillary pressure are shown in Fig. 3.}
\end{align*}
\]
Fig. 3—Relative permeabilities and capillary pressure used to construct the analytical solution shown. As expected, the curves show water-wet characteristics for (top) Ketton limestone and (bottom) diatomite.

There is one important difference between \( Q_w(t) \) in the Buckley-Leverett solution and the \( q_w(t) \) we have here. In the Buckley-Leverett theory, \( Q_w(t) \) is determined from the imposed injection rate. For the case of capillary-controlled flow, however, \( C \) and consequently \( q_w(t) \) cannot be imposed, but are determined by the rock’s ability to spontaneously imbibe water; thus, \( C \) is a parameter that describes an intrinsic property of the rock/fluid pair. In what follows, we will use the following relation from Eq. 11:

\[
\frac{d\omega}{dS_w} = \frac{2C}{\phi} F'(S_w). \tag{13}
\]

We now rewrite the derivatives of the water saturation:

\[
\frac{\partial S_w}{\partial t} = \frac{\partial S_w}{\partial \omega} \frac{\partial \omega}{\partial t} = -\omega \cdot \frac{1}{2\omega} \frac{dS_w}{d\omega}.
\]

\[
\frac{\partial S_w}{\partial x} = \frac{\partial S_w}{\partial \omega} \frac{\partial \omega}{\partial x} = \frac{1}{\sqrt{\frac{1}{\eta}}} \frac{dS_w}{d\omega}. \tag{14}
\]

By use of Eq. 14, the saturation equation, Eq. 6, can be rewritten as an ordinary-differential equation that is much simpler to solve because we only have to deal with total derivatives:

\[
\omega \frac{dS_w}{d\omega} + 2 \frac{d}{d\omega} \left[ D(S_w) \frac{dS_w}{d\omega} \right] = 0 \quad \text{(countercurrent case)}
\]

\[
\omega \frac{dS_w}{d\omega} + 2 \frac{d}{d\omega} \left[ D(S_w) \frac{dS_w}{d\omega} - 2Cf_w(S_w) \frac{dS_w}{d\omega} \right] = 0 \quad \text{(cocurrent case).} \tag{15}
\]

where for the cocurrent case we have used Eq. 10 to substitute for \( q_w \).

Eq. 14 can also be used to check that \( F \) truly is the fractional-flow function and satisfies Eq. 12. By inserting \( q_w(S_w) = q_w(S_w) F(S_w) \) from Eq. 5 into the original conservation equation, Eq. 6, we again obtain Eq. 11, which confirms the hypothesis in Eq. 12.

To remove the second-order derivatives, from Eq. 15, we integrate once

\[
\int \omega dS_w = \frac{2D}{\phi} \frac{dS_w}{d\omega} \quad \text{(countercurrent case),}
\]

\[
\int \omega - 2Cf_w(S_w) dS_w = -\frac{2D}{\phi} ES_w \quad \text{(cocurrent case),}
\]

\[
\ldots \ldots \ldots \ldots \ldots \ldots \ldots \ldots \ldots \ldots \ldots \ldots \ldots \quad \text{Eq. 16}
\]

where the integration constant is zero because \( D(S_{wmax}) = 0 \), and because \( F \) is the fractional-flow function \( F(S_{wmax}) = 0 \). We now substitute \( \omega dS_w = 2C/\phi dF \) and \( dS_w/d\omega = \phi(2CF)(S_w) \) from Eq. 13. This leads to a second-order ordinary-differential equation for \( F \):

\[
FF'' = -\frac{\phi}{2CF} D \quad \text{(countercurrent case),}
\]

\[
(F - f_w)F'' = -\frac{\phi}{2CF} D \quad \text{(cocurrent case).} \tag{17}
\]

Eq. 17 is the main mathematical result of this paper. From Eq. 17, we can determine \( F \) and \( C \), and once they are known, \( F' \) and \( C \) can be inserted into Eq. 11 to obtain the saturation profile. To determine \( F \), \( F' \), and \( C \) from Eq. 17, two possibilities exist. Formally, to obtain a closed-form expression for \( F \) and \( C \), Eq. 18 can
be integrated twice. This then leads to \( F \) and \( C \) being defined implicitly through two integral equations. McWhorter and Sunada (1990), Schmid et al. (2011), and Bjørnar and Mathiás (2013) solved these integral equations to obtain solutions; instead, we propose a simple backward-differencing scheme to solve the equations numerically.

The resulting algorithm is simple and can readily be used in a spreadsheet. This solution procedure is described in the next section. Before proceeding, however, we will briefly summarize the two important assumptions in this analysis. We assume incompressible 1D flow in an infinite medium. The solution is valid for a finite system until the leading edge of the saturation reaches the boundary. The time \( t^* \) when the solutions stop being valid in a finite matrix block of length \( L \) can be obtained from setting \( x(S_w,t^*) = L \), which yields

\[
  t^* = \left[ \frac{Lq}{2CF(S_{wi})} \right]^2. \tag{18}
\]

The time until \( t^* \) is called “early-time imbibition,” up to which our solutions are valid.

The second assumption we make is that the inlet-flow rate declines as \( 1/\sqrt{t} \) (Eq. 10). This can be rigorously proved as the correct boundary condition for SI controlled by capillary forces only (Schmid et al. 2011). However, when gravitational effects become significant, the solutions are no longer accurate, because the flow rate will either be enhanced (for instance, if the denser wetting fluid is moving downward) or suppressed (the denser wetting phase is moving upward).

### A Simple Way to Obtain \( F(S_w) \) and \( C \) Numerically

In this subsection, we describe how to obtain an approximation of arbitrary accuracy for \( F \) and \( C \) from Eq. 17. We determine \( F^i \) from a backward-differencing approximation, and then iteratively determine \( F(S_w) \) at a finite number \( n \) of saturation points \( S_{wi} \), \( j = 1, \ldots, n \) that are distributed evenly with a distance \( \Delta S_w \), between the residual and the maximum water saturation. Because \( F \) also depends on the unknown constant \( C \), we also need to iterate on \( C \); therefore, we start with a guessed value of \( C \) and continue changing it until \( F(S_w) \) converges to the correct solution. We assess convergence by testing if our computed value of \( F(S_w - S_{wi}) = 0 \).

The backward-difference scheme yields

\[
  F^i(S_w) \approx F\left(S_w + 2\Delta S_w\right) - 2F(S_w + \Delta S_w) + F(S_w). \tag{19}
\]

We now replace \( F^i \) in Eq. 17 with the approximation and rearrange. For the cocurrent case, this leads to

\[
  F(S_w) = \left[ F(S_w + 2\Delta S_w) - 0.5F(S_w + 2\Delta S_w) \right] + \sqrt{\frac{F(S_w + \Delta S_w) - 0.5F(S_w + 2\Delta S_w)}{2\Delta S_w} + \left( \frac{\phi}{2CF} \right)D(S_w)} \Delta S_w^{-1}. \tag{20}
\]

and for the cocurrent case, we have

\[
  F(S_w) = \left[ F(S_w + 2\Delta S_w) - 0.5F(S_w + 2\Delta S_w) + 0.5f_w(S_w) \right] + \sqrt{0.25\left[F(S_w + 2\Delta S_w) - 2f_w(S_w + \Delta S_w) - f_w(S_w)^2\right] - \left\{ f_w(S_w)^2 \right\}} + \sqrt{\frac{F(S_w + 2\Delta S_w) - 2f_w(S_w + \Delta S_w) + f_w(S_w)}{2CF}D(S_w)} \Delta S_w^{-1}. \tag{21}
\]

We therefore can find an approximation of \( F(S_w) \) if we know \( F(S_w + 2\Delta S_w) \) and \( F(S_w + 2\Delta S_w) \) for some starting value \( F(S_w) \). As the two starting values, we take \( F(S_{w,\max}) = 1 \) and \( F(S_{w,\max} - 2\Delta S_w) \).

The latter can be obtained by recalling that \( F(S_{w,\max}) \) can be approximated by a Taylor series:

\[
  F(S_{w,\max} - \Delta S_w) = F(S_{w,\max}) - \Delta S_w F'(S_{w,\max}) = 1, \tag{22}
\]

where we used the assumption that the maximum saturation is not transported into the medium, and hence \( F'(S_{w,\max}) = 0 \).

For the second criterion, we simply use that the fact because of mass balance, the integrated saturation curve must equal the total pore volume imbibed; that is,

\[
  \int_{S_{wi}}^{S_{w,\max}} x(S_w, t) ds = \frac{Q_w(t)}{\phi} = 2\sqrt{t}. \tag{24}
\]

We approximate the integral and rearrange Eq. 24 by use of Eqs. 9 and 11 to give

\[
  \sum_{j=1}^{n} F(S_w, i) \cdot \Delta S_w \approx \frac{Q_w(i)}{\phi} = 2\sqrt{i}. \tag{25}
\]

To determine \( C \), we start with some value \( \hat{C} \) and continue to change \( C \) until Eqs. 23 and 25 are met. The calculation procedure is fairly simple and has been incorporated into a spreadsheet that can be downloaded. The numerical procedure and the use of the spreadsheet are illustrated in the following examples.

### Examples: Water-Wet and Mixed-Wet Systems

In the following, three examples are given to illustrate the numerical procedure described previously. They show how to obtain the analytical solution for strongly and weakly water-wet and mixed-wet systems, and also demonstrate how the analytical solutions predict key physical features of water-wet and mixed-wet behavior during SI. The latter is of particular importance because the majority of fractured reservoirs are mixed-wet (Anderson 1987).

The input parameters used in the computations are listed in Table 1. In the following, we use power laws for the relative permeability and the capillary pressure curves. Because the analytical solution does not require specific functional forms, other models such as the Brooks-Corey or Van Genuchten (Helming 1997) could also be used (McWhorter and Sunada 1990; Schmid et al. 2011; Bjørnar and Mathiás 2013). The water and oil relative permeabilities are described by

\[
  k_{rw}(S_w) = k_{rw,\max} \left( \frac{S_w - S_{wi}}{1 - S_{wi} - S_{wi}} \right)^{\beta_r},
\]

\[
  k_{ro}(S_w) = k_{ro,\max} \left( \frac{1 - S_w - S_{wi}}{1 - S_{wi} - S_{wi}} \right)^{\beta_o}, \tag{26}
\]

where \( k_{rw,\max}, \alpha = w, o \) is the maximum relative permeability of water or oil, and \( \beta_r \) is the exponent of the relative permeability.

For the capillary pressure curve, we need to distinguish between the water-wet and mixed-wet cases. For the water-wet system, the \( p_i \) curve is positive over the whole saturation range, and SI only stops once the water saturation has reached \( S_w = 1 - S_{oi} \). The capillary pressure curve is described by

\[
  p_i(S_w) = \frac{S_w - S_{wi}}{1 - S_{wi} - S_{wi}} \tag{27}
\]
For a capillary entry pressure $p_{c, entry}$ and an exponent $\beta_{pc}$.

For a mixed-wet rock, SI stops at the saturation $S^*$ where $p_c$ becomes zero. In this case, the final recovery is governed by $S_w$, and not by $S_{or} = 1 - S_{ow}$. Here, we also use Eq. 27 for $p_c$, but then Eq. 27 only describes the positive part of the capillary pressure curve that is relevant for SI. The relative permeabilities and capillary pressures we use are shown in Fig. 4, while the corresponding capillary-fractional flow $F$ is shown in Fig. 5.

**Water-Wet Cases.** The relative permeabilities and the capillary pressure curve for the water-wet case were chosen such that they show the basic features of a water-wet system. The capillary pressure curve is larger than zero over the whole saturation range; the endpoint value of $k_{ow,max}$ and $k_{ow}$ itself are small, and the residual saturation $S_{or}$ is relatively large. Because SI can take place over the whole saturation range, the correct boundary value for the water saturation is $S_{w, max} = 1 - S_{or}$.

If $C$ and $F$ are known (Fig. 5), we can determine the $S_w$ profile that is shown in Fig. 6. The profile shows the diffusive, smooth behavior typical for SI processes. Notice also that $F$, contrary to its viscous counterpart $f_w$, has no inflection point (Fig. 5). Therefore, the saturation profile cannot develop a shock front. In this sense, this makes the SI solution easier to construct than the Buckley-Leverett problem.

From $C$, we can also determine how fast the rock can imbibite water and expel the oil; Fig. 7 shows the percentage of total recoverable oil over time from a core of 1-m length. $Q_w$ was calculated from Eq. 10 up to the point when the imbibition front reaches the end of the core at time $t^*$. We will now compare the behavior of cocurrent and countercurrent SI. We assume that in both cases we can use the same relative permeabilities, even though the few measurements in the

<table>
<thead>
<tr>
<th>Input Parameters</th>
<th>Strongly Water-Wet Case</th>
<th>Weakly Water-Wet Case</th>
<th>Mixed-Wet Case</th>
<th>Experiment (Ketton)</th>
<th>Experiment (Diatomite)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Initial water saturation</td>
<td>0.2</td>
<td>0.2</td>
<td>0.2</td>
<td>0</td>
<td>0</td>
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<tr>
<td>Residual oil/air saturation</td>
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<td>0.1</td>
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<td>0.02</td>
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<tr>
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<td>3</td>
<td>3</td>
<td>10</td>
<td>10</td>
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<tr>
<td>$k_{ow,max}$</td>
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<td>0.5</td>
<td>0.6</td>
<td>0.05</td>
<td>0.3</td>
</tr>
<tr>
<td>Exponent $\beta_{ow}$</td>
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<td>1.5</td>
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<tr>
<td>$k_{ow,min}$</td>
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<td>0.8</td>
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<td>0.95</td>
</tr>
<tr>
<td>$p_{c, entry}$ (Pa)</td>
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<td>12 000</td>
<td>12 000</td>
<td>35 000</td>
<td>665 000</td>
</tr>
<tr>
<td>Exponent $\beta_{pc}$</td>
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<td>-0.3</td>
<td>-0.2</td>
<td>-1.3</td>
<td>-1.2</td>
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<tr>
<td>Maximum capillary pressure (Pa)</td>
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<td>400 000</td>
<td>400 000</td>
<td>400 000</td>
<td>2 500 000</td>
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<td>$S^*$</td>
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<td>0.32</td>
<td>0.72</td>
<td>0.98</td>
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<tr>
<td>$\mu_w$ (Pa s)</td>
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<tr>
<td>$\mu_o$ (Pa s)</td>
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<td>0.001</td>
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<td>$10^{-6}$</td>
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<tr>
<td>$K$ (md)</td>
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<td>$\phi$</td>
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<td>0.2</td>
<td>0.2</td>
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<tr>
<td>$\sigma_{ow}$ (N/m)</td>
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<td>0.05</td>
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<td>$C_{cocurrent}$ ($m^3/m^3$)</td>
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<td>2.768$x10^{-5}$</td>
<td>0.0002</td>
<td>-</td>
</tr>
<tr>
<td>$C_{countercurrent}$ ($m^3/m^3$)</td>
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<td>2.733$x10^{-5}$</td>
<td>-</td>
<td>0.000217</td>
</tr>
</tbody>
</table>

Table 1—Input parameters for calculating the analytical solutions for strongly water-wet, weakly water-wet, and mixed-wet examples and for the cocurrent experiment. For the experiments, the nonwetting phase was air, and the rocks were Ketton carbonate and diatomite.

\[
p_c(S_w) = p_{c, entry} \left[ \frac{(S_w - S_{ow})}{(S_w^* - S_{ow})} \right]^{\beta_{pc}} - p_{c, entry}. \quad \ldots \ldots \ldots \ldots \ldots \ldots \ldots \ldots (27)
\]

for a capillary entry pressure $p_{c, entry}$ and an exponent $\beta_{pc}$.

Fig. 4—(a) Relative permeabilities and (b) the positive part of the capillary pressure curve plotted vs. $S_w$ for the strongly water-wet (sww), weakly water-wet (www), and mixed-wet (mw) examples presented in this paper. SI only takes place for positive capillary pressure and stops at $S_w$, which for a mixed-wet system is lower than $1 - S_{or}$.
For mixed-wet systems, SI stops at the typical of mixed-wet systems determine the size of interplay of relative permeabilities and capillary pressure curves. We now consider mixed-wet rock. The mixed-wet example are chosen such that they represent key features of mixed-wet systems (Behbahani and Blunt 2005; Ryzanov et al. 2009, 2010). The more oil-wetting the system is, the smaller the water saturation at which the capillary pressure goes negative, and the smaller the amount of oil that can be produced by SI.

Fig. 6 shows the comparison of the saturation profiles for the three wettabilities. The coherent profiles are again more piston-like than those for countercurrent SI, but as the sample becomes mixed-wet, the difference is small. For the mixed-wet system, the viscous-coupling term becomes negligible because $f_v(S_w)$ is very small in the relevant low-water-saturation range (Fig. 5).

Fig. 7 shows that the percentage of recoverable oil over the recovery rate is much lower for the mixed-wet case because imbibition is impeded by the low water relative permeability at low saturation and relatively low capillary pressure, compared with the water-wet cases. This dramatic increase in imbibition time for mixed-wet rocks has been observed experimentally (Zhou et al. 2000).

**Mixed-Wet Example.** We now consider mixed-wet rock. The interplay of relative permeabilities and capillary pressure curves typical of mixed-wet systems determine the size of $C$ and hence production rate (Eq. 10). For mixed-wet systems, SI stops at the saturation $S_{w,max}$ where $p_r$ becomes zero. Therefore, the boundary value for the water saturation $S_{w,max}$ has to be set to this saturation.

The relative permeabilities and capillary pressure curves for the mixed-wet example are chosen such that they represent key
Application: Matching to Experimental Data and a Scaling Group for Field-Scale Simulation

Analysis of Experiments. Fig. 2 shows experimental imbibition profiles scaled as a function of \( \sqrt{t} \). We can match the measurements to an analytical solution by use of the approach shown in this paper through adjusting the relative permeabilities and capillary pressure. The imbibition profile depends on three functions: two relative permeabilities and the capillary pressure. Hence, we cannot determine all three from this experiment alone. If we had measured relative permeabilities (by use of the classical Buckley-Leverett analysis), then the capillary pressure can be determined. In contrast, we could find the water relative permeability if the capillary pressure and oil relative permeability (from a centrifuge experiment, as an example) were known. Overall, the SI profile is a useful complement to other core measurements and can be used to determine multiphase-flow properties.

With an imbibition-saturation profile measured as a function of \( C \) (Eq. 11), we can integrate it to find \( F \) and differentiate it to find \( F' \) within the constant \( C \) (Eq. 13). \( C \) itself is found from the amount of water imbibed (Eq. 10). Then, from Eq. 7, we can determine the capillary dispersion \( D(S_w) \) directly. This is a more rigorous and direct approach than that presented by authors such as Zhou et al. (2002), who determined relative permeabilities and capillary pressure from imbibition measurements (such as those shown in Fig. 3) by use of a numerical approach. That is more cumbersome, is prone to numerical error, and did not achieve an exact match.

We adjusted the parameters shown in Table 1 to find a match to the experimental profiles for the SI experiments shown in Fig. 2. The value of \( C \) in this case is not obtained iteratively, as in the examples discussed previously, but is also determined experimentally from the relation between \( Q_c(t) \) and \( C \) (Eq. 10), fitting \( C \) to the measured water volume imbibed over time. The agreement between the calculated \( S_w \) profile and the measured one demonstrates that the theoretical predictions do indeed yield saturation profiles similar to those measured experimentally. The slight discrepancies between the experiment and the analytical solutions are likely because of heterogeneity in the rock samples, leading to a nonuniform imbibition profile.

Fig. 3 shows the relative permeabilities and capillary pressure used to construct the analytical solution. In the experiments with water imbibed into dry rock, we expect the system to be strongly water-wet. This is indeed what is observed with a low water relative permeability with a residual air saturation of 0.28 for Ketton limestone. The diatomite used by Zhou et al. (2002) showed weaker water-wet behavior compared with Ketton, with a higher water relative permeability and a low residual saturation.

Scaling Group for Field-Scale Simulation. The second application of this work is to determine scaling groups for use in larger-scale simulation, such as in dual-porosity modeling of fractured reservoirs. Schmid and Geiger (2012) introduced the nondimensional time for SI as the total volume imbibed divided by pore volume:

\[
t_d = \left( \frac{Q_c(t)}{\phi L} \right)^2 = \left( \frac{2C}{\phi L} \right)^2.
\]

Note that for SI the expression in the brackets needs to be squared because, as we have seen previously, \( Q_c(t) \) is proportional to \( \sqrt{t} \). Schmid and Geiger (2012, 2013) have shown that this scaling group can be used to reduce experimental measurements of overall recovery as a function of time onto one universal curve, allowing the average SI behavior to be captured elegantly and accurately.

This scaling group can then be used in dual-porosity models of field-scale flow in fractured reservoirs, work that is under development (Di Donato et al. 2007; Geiger et al. 2013; Maier and Geiger 2013; Maier et al. 2013).

Summary and Conclusions

In this paper, the general analytical solutions for SI were derived following McWhorter and Sunada (1990) and Schmid et al. (2011). These solutions allow for important insights into the physics of SI. These insights enable the following:

1. The construction of solutions by use of a simple iterative approach on a spreadsheet.
2. The quantification of wettability effects on imbibition rate, which exhibits a considerable slowing of recovery in mixed-wet systems because of the combined effects of low water relative permeability and capillary pressure.
3. The prediction of significant differences between cocurrent and countercurrent SI, in line with experimental results. Cocurrent imbibition is more piston-like and leads to higher production rates than countercurrent SI. However, for mixed-wet systems, this effect becomes weak and the difference between cocurrent and countercurrent SI becomes negligible.
4. The demonstration of how experimental measurements of saturation profiles in imbibition can be used to obtain or constrain relative permeabilities and capillary pressures by matching experimental measurements to the analytical solutions.

Nomenclature

\[
\begin{align*}
C & = \text{constant that quantifies the rate of SI, m}/\sqrt{s} \\
D & = \text{capillary dispersion, m}^2 \\
f & = \text{fractional flow for viscous-controlled flow} \\
F & = \text{fractional flow for capillary-controlled flow} \\
k_r & = \text{relative permeability} \\
K & = \text{absolute permeability, m}^2 \\
L & = \text{length of the core, m} \\
p_c & = \text{capillary pressure, Pa} \\
P & = \text{fluid pressure, Pa} \\
Q_c & = \text{cumulative volume imbibed/injected, m}^3 \\
R & = \text{recovery, m}^3 \\
R_u & = \text{ultimate recovery, m}^3 \\
S & = \text{saturation} \\
t & = \text{time, seconds} \\
t_d & = \text{dimensionless time} \\
\beta & = \text{exponent for the relative permeabilities and capillary pressure} \\
\lambda & = \text{mobility, 1/(Pa-s)} \\
\mu & = \text{fluid viscosity, Pa/s} \\
\sigma & = \text{interfacial tension, N/m} \\
\phi & = \text{porosity}
\end{align*}
\]

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References


